Dissertation Introduction

Notes:

* Explain Moho and how it is studied and list some people or models that have either tried to make local or global models.
* More detail in how Moho model is calculated by removing everything else, and uncertainties in estimates arise from unmodelled masses. Add in underplating from Mariani 2013 in Parana basin.
* Swzillus has uncertainties, in interpolation of seismic results, look a paper.
* Go into case specific, lots of models have estimates but few have uncertainties in these estimates.
* How model is calculated from gravity using seismic constraints.
* Aim is to find uncertainties by using cross validation to estimate the difference between gravity and seismic and see how good the gravity data is where there is no seismic. Also to compare it to e.g. Amazon Rainforest, where seismic surveys are almost impossible to carry out.
* Cross validation builds upon Uieda method as it involves a training and testing set which are different from one another.

Introduction:

The Mohorovičić discontinuity or Moho for short is the physical boundary signified by the change in many properties such as mineralogy, density, and temperature among other things, but it is mainly known as the change from the crust to the upper mantle. Ever since the discovery of this boundary observed through a significant change in seismic p-wave velocity on either side of the moho by seismologist Andrija Mohorovičić in 1909 there have been many people using multiple methods to try and quantify the depths to this boundary. Some estimates of the Moho discontinuity include Laske et al. (2013), Assumpção et al (2013) and Reguzzoni and Sampietro (2015), all of which have slightly differing models depending on types of methods and data. Geophysical techniques such as seismology and gravimetry have been utilized to estimate the depth of the discontinuity and topography over a local to a global scale. This issue is known as a geophysical inverse problem and uses data in the form of gravity, seismic or another method to determine the depth to the Moho over a certain area and produce a model. In this paper the moho model will be derived from gravitational data and removing the local gravity disturbances to get the regional field that is almost entirely based off the depth to the Crust-Mantle boundary with the help of other parameters that need to be constrained including the density, reference Moho, and regularisation parameter. The problem with using gravitational data to model the Moho is that gravity values calculated to be the regional field does not only include the effect of the discontinuity but also the effects of unmodelled masses in the Earth’s crust that have not been removed when taking away all other effects that contribute to the strength of the raw gravity data. These unmodelled masses lead to the emergence of uncertainties in the model created, however with these uncertainties being of an unknown magnitude then it is difficult to quantify and correct for these unknown masses as the location, size, density, and number are impossible to determine. This is not just the case in the Uieda & Barbosa (2017), but almost all gravitationally derived Moho models suffer the same fate including van der Meijde et al. (2013), Tugume et al. (2013), Gimenez et al. (2019), and Reguzzoni and Sampietro (2015). All of these models produced are generally accurate representations of the Moho surface however none of these include errors or uncertainty values in to coincide with their models. Although these error estimates are often not seen in seismologically derived Moho models either, one of the only recognisable papers that tries to calculate uncertainty of their estimates is Szwillus et al. (2019) who obtain uncertainties by interpolating Moho depth even then these uncertainties calculated are larger than the errors coming from the P-wave velocity.

In this paper the aim is to find uncertainties in a gravitationally calculated Moho model by using cross validation with seismic constraints. With the hopes of finding the difference between a model calculated from gravity data and seismic point estimates to see how good the gravitational estimates of the depth to the moho are where there are not any seismic points that the model is able to constrain to. Continents such as Europe and North America have extensive seismological surveys that span most of the area so it would be considered pointless to see how well this method of uncertainty estimation works as there are not substantial areas of land without seismic data. This is why South America has been chosen as most of the surveyed areas by either reflection or refraction are situated along the coast with few based towards the centre of the continent mainly due to the difficultly of undertaking a survey as a result of the magnitude and density of vegetation in the Amazon rainforest. However, South America is also limited with seismic data due to lack of financial funding as survey over a large scale may not be economically viable for countries or companies that are interested in carrying out one. So with the vast area of the Amazon having little to no seismic points to constrain the gravitationally computed Moho discontinuity model it will be a good test to see how good the model is when there are no seismic point estimates to compare to.

Therefore this paper will use the Moho model of South America created by Uieda & Barbosa (2017) and implement a cross validation approach to calculate the uncertainties through the average differences between gravitational estimates and seismic point estimates of the crustal thickness or depth to Moho. The method builds upon the cross validation originally used in the paper by randomly selecting a training and testing set and calculating not only the uncertainty but how many seismic point constraints are needed to accurately quantify the size of this error. In addition to using cross validation, in the synthetic models created of South America the effect of trying to model a previously unknown mass will be tackled by adding in underplating in the Parana basin addressed in Mariani et al. (2013).